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# **A Well-balanced Finite Volume Scheme for** Shallow Water Equations with Porosity - Nº22 Minh H. Le<sup>1</sup>, Virgile Dubos<sup>3</sup>, Marina Oukacine<sup>1</sup> and Nicole Goutal<sup>1 2</sup> <sup>1</sup>Saint-Venant Hydraulics Laboratory (LHSV), <sup>2</sup>Electricity of France (EDF) R&D, Chatou, <sup>3</sup>Sorbonne University, Inria, LJLL, ANGE

| Framework  | Numerical experiment   |
|--|--|
| Modelling emergent and rigid vegetation in open-channel with a single porosity-based | <sup>4.44</sup> Reference solution well-known test case of SW model. |

- shallow water model (SP) [1].
- New Godunov-type, finite volume, well-balanced and shock capturing scheme for SP.
- Comparison with other methods in the literature [2, 3] and first application.

## Single porosity shallow water model

$$\partial_{t}W + \partial_{x}F + \partial_{y}G = S_{b} + S_{\phi} + S_{\tau} \qquad (1$$
with  $W = \phi \begin{pmatrix} h \\ hu \\ hu \\ hv \end{pmatrix}, F = \phi \begin{pmatrix} hu \\ hu^{2} + \frac{g}{2}h^{2} \\ huv \end{pmatrix}, G = \phi \begin{pmatrix} hv \\ huv \\ hv^{2} + \frac{g}{2}h^{2} \end{pmatrix},$ 

$$S_{b} = \phi \begin{pmatrix} 0 \\ -gh\partial_{x}b \\ -gh\partial_{y}b \end{pmatrix}, S_{\phi} = \begin{pmatrix} 0 \\ \frac{g}{2}h^{2}\partial_{x}\phi \\ \frac{g}{2}h^{2}\partial_{y}\phi \end{pmatrix}, S_{\tau} = -\begin{pmatrix} 0 \\ g\phi h\frac{\eta^{2}|U|}{h^{4/3}}u + \frac{1}{2}\frac{aC_{d}h|U|}{\phi}u \\ g\phi h\frac{\eta^{2}|U|}{h^{4/3}}v + \frac{1}{2}\frac{aC_{d}h|U|}{\phi}v \end{pmatrix}.$$
where
$$\begin{cases} h : \text{ depth of water,} \\ b : \text{ bed elevation,} \\ \eta : \text{Manning's coefficient,} \\ C_{d} : \text{ drag coefficient,} \end{cases} \begin{cases} u, v : \text{ depth-averaged horizontal velocities,} \\ g : acceleration due to gravity, \\ D : effective diameter of vegetation, \\ a = \frac{1-\phi}{D}\mu} : frontal area of vegetation. \end{cases}$$

 $\pi D/4$ 

## Numerical scheme

 $\phi$  :

## <u>Well-balanced scheme</u> : we solve (1) in two following steps:

porosity,



Figure: Steady sub-critical flow over a bump.

- test case from [3], exact solution known.
- complex structure: 1-rarefaction wave, stationary discontinuity, 1-rarefaction wave then 2-shock wave.
- hydrostatic approximation: wrong solution unlike Bernoulli resolution.
- PorAs scheme can have difficulties to estimate shock waves speed.

## Bernoulli relation in (6): $\frac{u_L^{*2}}{2g} + h_L^* + b_L = \frac{u_R^{*2}}{2g} + h_R^* + b_R.$ hydrostatic approximation yields:

SēMA

- $h_{I}^{*} + b_{L} = h_{R}^{*} + b_{R}.$
- ▶ as efficient as the PorAS scheme [2] for this type of regular solution.



Figure: Dambreak with large porosity discontinuity.

Longitudinal transition from meadow to wood in a laboratory flume [4]:

Convection:Friction and Drag:
$$\begin{cases} \partial_t W + \partial_x F + \partial_y G = S_b + S_{\phi}, \\ W(x, 0) = W^n. \end{cases}$$
(2) $\begin{cases} \partial_t W = S_{\tau}, \\ W(x, 0) = W^{n+1/2}. \end{cases}$ (3) $\underline{FV}$  scheme for (2):  $W_j^{n+1/2} = W_j^n - \frac{\Delta t}{|C_j|} \sum_{\Gamma_{jk} \subset \partial C_j} |\Gamma_{jk}| \mathcal{F}(W_j^n, W_k^n; b_j, b_k).$ Rotational invariance of the flux  $\Rightarrow$  numerical flux derived from the 1D system (4).(3)

problem with  $S_X(VV, D) =$  $\begin{array}{ll} W(x,0) = \begin{cases} W_L & \text{if } x < 0, \\ W_R & \text{if } x > 0. \end{cases} \begin{array}{ll} (0, \frac{g}{2}h^2\partial_x\phi - g\phi h\partial_x b, 0) \text{ and } W_{L,R} \text{ the pro-} \\ \text{jected states } W_{j,k}^n \text{ on } \Gamma_{jk}. \end{cases} \end{array}$ 

Approximation by simple-solver  $W_{\mathcal{R}}(x/t)$  of (4):



First order Godunov-type scheme :  $W_{j}^{n+1/2} = W_{j}^{n} - \frac{\Delta t}{\Delta x} \left( F_{j+1/2}^{L} - F_{j-1/2}^{R} \right)$  $\begin{cases} F^L = F(W_L) + \lambda_L(W_L^* - W_L) + \lambda_-^*(W^* - W_L^*), \\ F^R = F(W_R) - \lambda_R(W_R - W_R^*) - \lambda_+^*(W_R^* - W^*). \end{cases}$ 

• (7) and 2nd invariant in (6) lead to:  $h_L^* = rac{(\phi h)^{HLL} + lpha \phi_R (b_R - b_L)}{lpha \phi_R + (1 - lpha) \phi_L},$ References  $h_R^* = rac{(\phi h)^{HLL} - (1-lpha)\phi_L(b_R - b_L)}{lpha \phi_R + (1-lpha)\phi_L}.$ • (5) and 3rd invariant (6) yield:  $\lambda^* = \frac{\phi_L h_L (u_L - \lambda_L) + \phi_R h_R (u_R - \lambda_R)}{2(\phi h)^*} + \frac{\lambda_L \phi_L h_L^* + \lambda_R \phi_R h_R^*}{2(\phi h)^*}$  $\Delta x \ \overline{S_x} = g\left(\frac{h_L h_R}{2}(\phi_R - \phi_L)\right)$  $-g\left(\frac{\phi_L h_L + \phi_R h_R}{2}(b_R - b_L)\right)$ 



Figure: Velocity field and unit discharge in wood region: SW model (top) and SP model (bottom).

- emergent, rigid and circular vegetation distributed in staggered rows.
- expensive computation cost with SW.
- SP capture macroscopic behaviours.

## **Conclusions & Forthcoming Research**

- augmented HLLC scheme well-balanced, positivity preserving and shock capturing.
- further test cases have to be implemented.

$$(1)$$

Figure: Water depth around the transition.

- SP capture qualitative behaviour. ▶ agreement in wood region,  $C_d = 1.2$ . ▶ better compromise with  $C_d = 1$ .
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